

ORIGINAL ARTICLE

Influence of geogenic factors on microbial communities in metallogenic Australian soils

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Links between microbial community assemblages and geogenic factors were assessed in 187 soil samples collected from four metal-rich provinces across Australia. Field-fresh soils and soils incubated with soluble Au(III) complexes were analysed using three-domain multiplex-terminal restriction fragment length polymorphism, and phylogenetic (PhyloChip) and functional (GeoChip) microarrays. Geogenic factors of soils were determined using lithological-, geomorphological- and soil-mapping combined with analyses of 51 geochemical parameters. Microbial communities differed significantly between landforms, soil horizons, lithologies and also with the occurrence of underlying Au deposits. The strongest responses to these factors, and to amendment with soluble Au(III) complexes, was observed in bacterial communities. PhyloChip analyses revealed a greater abundance and diversity of *Alphaproteobacteria* (especially *Sphingomonas* spp.), and *Firmicutes* (*Bacillus* spp.) in Au-containing and Au(III)-amended soils. Analyses of potential function (GeoChip) revealed higher abundances of metal-resistance genes in metal-rich soils. For example, genes that hybridised with metal-resistance genes *copA*, *chrA* and *czcA* of a prevalent aurophilic bacterium, *Cupriavidus metallidurans* CH34, occurred only in auriferous soils. These data help establish key links between geogenic factors and the phylogeny and function within soil microbial communities. In particular, the landform, which is a crucial factor in determining soil geochemistry, strongly affected microbial community structures.

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Introduction

Identifying drivers of microbial community structures in soils is challenging. Difficulties arise because of the complexity of soil ecosystems, where large number of niches provide for high levels of species diversity and complex ecosystem

interactions (Fierer and Jackson, 2006). Environmental factors, for example, soil-type, vegetation, landuse and associated physicochemical parameters, for example, C–N–, water content and pH, have been linked to soil microbial structures and activities (Acosta-Martinez *et al.*, 2008; Lauber *et al.*, 2008; Wakelin *et al.*, 2008; Drenovsky *et al.*, 2010). In contrast, the importance of geogenic drivers, comprising geomorphological, geological and geochemical factors such as the landform, the underlying lithology and the presence of buried mineral deposits (expressed in overlying soils as elevated concentrations of mobile metals) are less well understood (Viles *et al.*, in press). As geogenic influences are likely to develop over extended ‘geologic’ periods of time, their effects may be

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masked by 'short term' changes in vegetation, landuse or anthropogenic influences (Viles *et al.*, in press). However, these are likely to be second order influences within ecosystems, which are predominantly driven by geogenic factors.

High concentrations of metals, for example, soils exposed to industrial pollution or mining, substantially alter microbial communities (Baker and Banfield 2003; Hu *et al.*, 2007; Kock and Schippers, 2008; Denef *et al.*, 2010). Similarly, amendment of soils with Cu, Pb, Zn and Cd strongly influences resident bacterial communities (Bååth, 1989; Abaye *et al.*, 2005; Wakelin *et al.*, 2010 a, b). In contrast, few studies have looked at natural systems in which physical and biogeochemical processes have formed 'enrichment haloes' of metals in soils overlying mineral deposits (Aspandiar *et al.*, 2008). One study of soils overlying a base metal (Cu, Pb and Zn) deposit in Western Australia has shown that the solubilisation, transport and deposition of metals are mediated by resident plant and microbial communities, in turn altering the structure of microbial communities in the metal-rich soils overlying the deposit (Wakelin *et al.*, 2012b). Other studies have shown that soil microbial communities mediate Au mobilisation, and that underlying Au deposits may influence microbial communities (Reith and McPhail, 2006; Reith and Rogers, 2008).

This is the first study to use a comprehensive set of molecular tools, that is, multiplex-terminal restriction fragment length polymorphism (M-TRFLP), and high-density phylogenetic (PhyloChip) and functional (GeoChip) microarrays, in combination with field mapping and geochemical analyses, to assess the influences of geogenic factors on soil microbial communities. Hence, we assess: (i) if soil microbial community structure and functioning potential (as interpreted from the occurrence of specific genes on GeoChip microarrays) are influenced by landform, lithology and buried Au deposits; (ii) how underlying geochemical parameters define these influences; and (iii) if environmentally relevant concentrations of Au affect microbial communities.

Materials and methods

Field sites and sampling

Minimal landuse sites overlying Au deposits in remote Australia were sampled. Here, soils have developed over thousands to millions of years under tectonically stable conditions. Auriferous and non-auriferous soils were collected from four sampling areas, that is, Old Pirate, Tomakin Park, Humpback and Wildcat; a description of environmental characteristics and references is given in Supplementary Table S1. Old Pirate is located in the Tanami region of the Northern Territory (Figure 1). Across the site soil-type, vegetation and underlying lithology are uniform, allowing for the influence of the landform, as the dominant geogenic factor, to be assessed

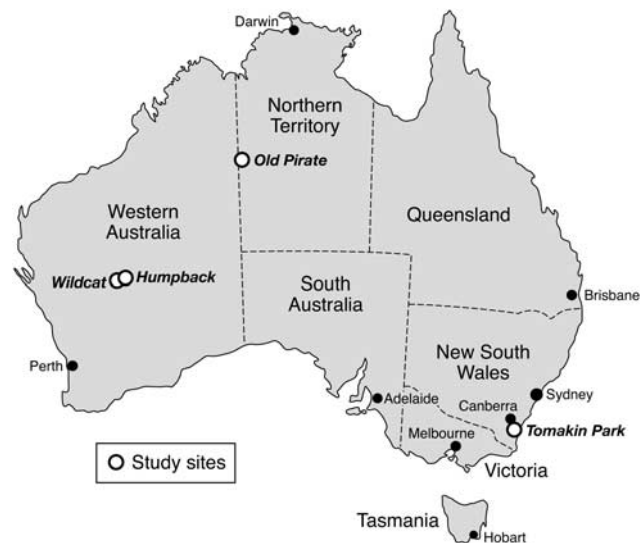


Figure 1 Map of Australia showing the Old Pirate, Tomakin Park, Humpback and Wildcat sites in the Northern Territory, New South Wales and Western Australia, respectively.

(Supplementary Table S1). Soils were collected in October 2009 from the A-horizon (at 0.1–0.15 m depth) along a nine kilometre traverse covering erosional, colluvial and alluvial landforms. The Tomakin Park site is located in New South Wales (Figure 1). The site has a uniform underlying lithology (phyllitic schist), landform (colluvium) and vegetation settings; hence the influence of soil horizons was assessed (Supplementary Table S1). Soils were collected in October 2007 along a 400-m traverse at depths of 0.03–0.05 and 0.15–0.2 m for Ah- and B-horizons, respectively. The (contiguous) Wildcat and Humpback sites are located in the Lawlers district of Western Australia (Figure 1). Samples were collected in June 2008 from the A-horizon (0.1–0.15 m) over an 8-km polygon covering different lithologies (mafic and granitic rocks) and landforms (erosional, colluvial, alluvial and depositional).

At each of the 187 sampling locations, 50-ml centrifuge tubes of soil were collected under field-sterile conditions. Landforms and soil types were classified using the Australian landform and soil classification systems, respectively (Isbell, 2002; Pain, 2008). Samples for DNA extraction were frozen on-site. Tubes stored at ambient temperature were used for geochemical analyses. At three locations, 10 kg of soil was collected for microcosm experiments. These soils were homogenised, sieved to <2 mm fraction, stored in sterile plastic bags and refrigerated at 4 °C.

Soil microcosms experiments

To assess the effect of soluble Au(III) complexes on community assemblages, microcosm experiments were conducted with Old Pirate, Wildcat and Humpback soils, following Reith and Rogers (2008).

Microcosms consisted of sterile 50 ml tubes containing 40 g d.w. (dry weight) soil. Soils were amended with 0, 50, 1000 and 100 000 ng Au(III)-chloride g⁻¹ d.w. soil and incubated aerobically at 90% water holding capacity in the dark. At 0, 2, 10 and 30 days, three replicates tubes were sampled from each experiment; DNA was extracted and analysed by M-TRFLP.

Assessment of community assemblages and functional potential

Nucleic acid extraction, quantification and quality control. For M-TRFLP, DNA was extracted in duplicate from 0.25 g of homogenised field and microcosm soils using the PowerSoil DNA Isolation kit (Mo Bio, USA). For microarray analyses, DNA was extracted in duplicate from 10 g of soil using the PowerMAX Soil Mega Prep DNA Isolation kit (Mo Bio). Duplicate extractions were combined, DNA was purified (phenol:chloroform), precipitated with isopropanol and quantified (Quant-iT; Invitrogen).

Multiplex-terminal restriction fragment length polymorphism. Bacterial, archaeal and fungal communities were characterised using M-TRFLP of the small subunit rRNA gene (Singh *et al.*, 2006). Primer sets and reagents are listed in Supplementary Table S2. Multiplex-PCR (PCRs, 25 µl) used Qiagen HotStar Taq chemistry and PCR consisted of 30 cycles of 95 °C for 30 s, 55 °C for 30 s, 72 °C for 60 s and a final extension step for 10 min at 72 °C (Singh *et al.*, 2006). Lengths of the final products are listed in Supplementary Table S2 and were verified via gel electrophoresis. PCR products were purified (WizardSV; Promega), and 100 ng digested with 20 U of *MspI*, *HaeIII* and *TaqI* for 3 h at 37 °C or 65 °C. Capillary separation of TRFs was conducted by the Australian Genome Research Facility. TRFs were scored in GeneMarker (SoftGenetics, USA) at a detection limit of 200 fluorescent units; TRFs differing ± 0.5 base pairs were binned together. Peak heights were used as a measure of abundance while richness was based on the number of TRFs obtained.

Similarity matrices were generated on square root transformed abundance data using the Bray–Curtis method (Bray and Curtis, 1957). Permutational multivariate analysis of variance (PERMANOVA) combined with canonical analysis of principal coordinates (CAP; Anderson *et al.*, 2008) were used to test the effects of geogenic factors. PERMANOVA analyses were conducted using partial sums of squares, on 9999 permutations of residuals under a reduced model. Associations between microbial and geochemical data sets were assessed using RELATE tests (permutation-derived Mantel-type testing). When significant ($P < 0.05$) relationships occurred, distance-based linear modelling was used to assess geochemical parameters best explaining the variability within the microbial data set. Geochemical parameters significant influencing

microbial structures are shown as vector overlays in CAP plots. Multivariate analysis was conducted using PRIMER6 and PERMANOVA+ (Primer-E Ltd, UK). Statistical routines are described in Clarke and Warwick (2001) and Anderson *et al.* (2008).

Microarray analysis (PhyloChip and GeoChip). The G2-PhyloChip microarray was used to characterise the bacterial community composition (Brodie *et al.*, 2006). Key samples were chosen that best represented influence by geogenic factors (CAP analyses on the geochemical data sets). Amplification of 16S rRNA genes, purification of products, labelling of DNA and hybridisation were conducted following Brodie *et al.*, (2006) using the reaction chemistry described in Wakelin *et al.* (2012b). Hybridised arrays were stained and washed on an Affymetrix fluidic station (Brodie *et al.*, 2006). Following scanning, data were processed following Brodie *et al.* (2006) and DeSantis *et al.* (2007). Data were imported into PhyloTrac for scoring of taxa (Schatz *et al.*, 2010). Operational taxonomic units were deemed detected by a positive fraction of probe-pair matches ≥ 0.9 . Taxa were ranked by relative abundance, and the 500 taxa with the highest intensities across all samples were aggregated into a data set, representing the dominant community. The taxonomical hierarchy for each taxon was determined and the distribution of phyla/classes plotted; taxa representing <1% of the total abundance were combined as ‘other’. Data were aggregated to phylum/class levels and log transformed. Similarity percentage analysis was used to identify groups that discriminate between auriferous and non-auriferous soils ($P < 0.1$; Clarke and Warwick, 2001).

Key samples from Old Pirate and Tomakin were analysed using GeoChip 3.0 (He *et al.*, 2010). Arrays were scanned by ScanArray Express Microarray Scanner (Perkin Elmer, USA) at 633 nm using a laser power of 90% and a photomultiplier tube gain of 75%. ImaGene v.6.0 (Biodiscovery, USA) was used to determine the quality of each spot. Raw data from ImaGene were submitted to Microarray Data Manager (<http://ieg.ou.edu/microarray/>; He *et al.*, 2010). In PRIMER6 data were log-transformed and hierarchically aggregated by gene name and category. Bray–Curtis similarity indices on log-transformed data were used to calculate distance matrices (Clarke, 1993). Similarity percentage analysis was used to identify genes and gene categories that discriminate between auriferous and non-auriferous soils ($P < 0.1$; Clarke and Warwick, 2001).

Geochemical characterisation

After homogenisation, 0.5 g of each soil was microwave digested in concentrated *aqua regia*. Concentrations of major metals were determined by inductively coupled plasma-optical emission spectrometry (Spectro ARCOS SOP, Germany);

minor- and trace metals and metalloids were determined by inductively coupled plasma-mass spectrometry (Agilent 7500ce, USA). Total C and N were determined by high-temperature combustion (Formacs analyser; Skalar Inc., USA); electrical conductivity (E.C.) and pH were measured in 1:5 soil to water extracts. Location coordinates, geogenic factors and geochemical data and detection limits are shown in Supplementary Table S3.

Data (except pH) were log transformed to remove skew. The data were normalised and similarity matrices based on Euclidean distances were generated. Physiochemical parameters were categorised into seven groups, including solution parameters (E.C. and pH) and six groups based on a modified Goldschmidt element classification (Goldschmidt, 1954; McQueen, 2008): Biophile- (C_{tot} , N_{tot} , P and S), calcophile- (Cu, Ga, Ge, Sn and Zn), Au-pathfinder- (Ag, As, Au, Bi, Mo, Pb, Se and W), lithophile- (Al, Ca, Cr, K, Mg, Na, Nb, Sc, Sr, Th, U and V), rare earth (Ce, Dy, Er, Eu, Gd, Ho, La, Lu, Nd, Pr, Sm, Tb, Tm, Y and Yb) and siderophile- (Fe, Co, Mn, Ni and Os) elements. PERMANOVA combined with CAP analyses were used to test the effects of underlying Au deposits, lithology, soil horizon and landform on geochemical parameters. Significant individual parameters are shown as vector overlays in CAP plots.

Results

Soil geochemistry

Geochemical data demonstrated highly significant

($P < 0.001$) associations between element groupings, their distribution and geogenic factors (Table 1). CAP analyses across sites, showed that the *a priori* classification of elements produced a highly significant differentiation ($P < 0.001$; Supplementary Figure S1). Distribution of elements was controlled by geogenic factors ($P < 0.05$; PERMANOVA), as reflected in $CV(\sqrt{\quad})$ values (Table 1; Anderson *et al.*, 2008). At Old Pirate, the concentration of elements in soils was strongly linked to the underlying Au deposit and landform ($P < 0.05$, Table 1), and a significant interactive effect between these parameters was observed (Table 1). In Tomakin, the element distributions differed with respect to the Au deposit and soil horizon ($P < 0.001$; Table 1). At Wildcat and Humpback the dominant factor controlling the element distribution was lithology ($P < 0.001$), with underlying Au deposit and landform displaying secondary influences (Table 1). CAP analyses revealed highly significant ($P = 0.001$) separation according to geogenic factors. Canonical correlations were high ($\delta 1 > 0.95$; $\delta 2 > 0.86$), and canonical axes (CAs) 1 and 2 accounted for $> 40\%$ and 10–20% of variation, respectively (Figures 2A, C and E).

M-TRFLP—fingerprinting of microbial communities in field soils

Significant links existed between geogenic factors and bacterial and fungal communities at all sites (Mantel coefficients 0.3; $P < 0.05$). At Old Pirate, the

Table 1 Results of PERMANOVA testing of the influence of the factors underlying Au deposit, soil horizon, landform and lithology on microbial communities and geochemical parameters at the study sites

Community structure	Bacteria ($\sqrt{\quad}$) CV ^f	Pseudo-F	P	Fungi ($\sqrt{\quad}$) CV	Pseudo-F	P	Archaea ($\sqrt{\quad}$) CV	Pseudo-F	P	Geochemistry ($\sqrt{\quad}$) CV	Pseudo-F	P
<i>Old Pirate</i>												
underlying D	5.1^s	1.5	0.09	6.5	1.3	0.1	−10.65	0.2	1.0	2.47	2.6	0.04
RL	8.0	2.6	0.0002	7.2	1.5	0.05	12.27	2.3	0.01	2.67	3.1	0.008
D × RL ^a	7.4	1.9	0.02	8.9	1.7	0.004	9.36	1.7	0.04	3.03	3.6	0.0006
Residual	23.1			39.2			40.49			6.93		
<i>Tomakin</i>												
Underlying D	18.5	4.6	0.0001	11.6	5.6	0.0001	10.8	1.6	0.04	3.67	7.3	0.0001
SH	13.6	2.9	0.0001	2.9	1.3	0.2	−2.7	1.0	0.5	3.01	5.2	0.002
D × SH ^b	7.8	1.3	0.07	−2.5	0.9	0.5	−11.4	0.7	0.9	2.45	2.4	0.06
Residual	41.7			23.3			59.9			6.31		
<i>Wildcat/Humpback</i>												
Underlying D	−7.8	0.7	0.8	5.0	1.5	0.09	24.5	1.6	0.2	2.54	2.8	0.03
L	10.6	1.8	0.03	15.8	3.8	0.0001	3.5	1.1	0.4	5.21	14.0	0.0001
RL	9.7	1.3	0.07	7.0	1.3	0.05	18.0	1.6	0.07	1.69	1.6	0.08
D × L ^c	5.7	1.1	0.4	4.0	1.1	0.4	10.3	1.0	0.5	1.41	1.2	0.3
D × RL	−11.0	0.7	0.9	4.3	1.1	0.3	12.5	1.2	0.3	1.23	1.2	0.3
L × RL ^d	3.3	1.0	0.4	11.0	1.2	0.1	23.4	1.4	0.1	2.49	6.2	0.07
D × L × RL ^e	−19.5	0.6	0.9	−14.3	0.8	0.7	Undetermined			1.40	1.1	0.3
Residual	42.4			40.0			54.7			5.65		

Abbreviations: D, deposit; L, lithology; PERMANOVA, permutational multivariate analysis of variance; RL, regolith landform; SH, soil horizon.

^aInteractive effect of underlying D and RL.

^bInteractive effect of underlying D and SH.

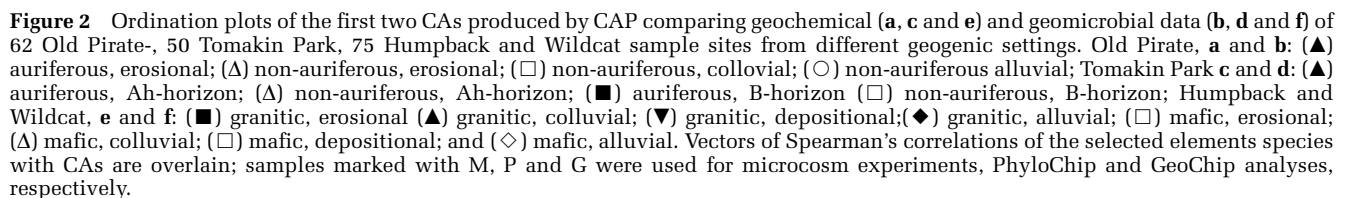
^cInteractive effect of underlying D and L.

^dInteractive effect of L and RL.

^eInteractive effect of underlying D, L and RL.

^f($\sqrt{\quad}$)CV is the square root of the component of variation, which is a data set dependent measure of the effect of size in units of the community dissimilarities (that is, increasing positive values); negative values indicate zero components (Anderson *et al.*, 2008).

^sSignificance level to assess the criteria was $P < 0.1$ (in bold).



bacterial community composition varied with landform ($\sqrt{CV}=8.0$; $P<0.001$) and underlying Au deposit ($\sqrt{CV}=5.1$; $P<0.1$), and these factors were interactive ($\sqrt{CV}=7.4$; $P=0.02$; Table 1). Similarly, soil fungal and archaeal communities varied across landforms, but not with the underlying Au deposit ($P<0.05$). An interactive effect ($P<0.05$) of landform and Au deposit was observed (Table 1). CAP showed that bacterial and fungal communities were strongly separated between auriferous and non-auriferous soils ($P<0.001$; Figure 2B; Supplementary Figure S2A). Canonical correlation coefficients were high ($\delta 1=0.91$; $\delta 2=0.70$), and CA1 and CA2 accounted for 34.2% and 12.0% of variation in the bacterial data set, respectively. To assess which geochemical parameters were linked to these variations, influences of each element group were evaluated using marginal testing in the DSTLM routine (Anderson *et al.*, 2008). Individually, all groups of parameters had a significant relationship with bacterial and fungal data sets. Solution parameters (pH and E.C.), and biophile-, lithophile- and siderophile major elements displayed the strongest relationships (Table 2). Minor elements were also capable of explaining a proportion of the spread, for example, Au-pathfinder elements explained 21.6% of variation in the bacterial data set (Table 2).

At Tomakin, the bacteria community varied with the underlying Au deposit ($\sqrt{CV}=18.5$; PERMANOVA $P<0.001$) and soil horizon ($\sqrt{CV}=13.6$; $P<0.001$). Fungal and archaeal communities varied with respect to the Au deposit ($P<0.05$), but not soil horizons (Table 1). Similarly, CAP analysis revealed a significant separation of bacterial and fungal communities ($P<0.001$; Figure 2d). Canonical correlations were high ($\delta 1=0.91$; $\delta 2=0.80$), and CA1 and CA2 accounted for 13.8% and 10.5% of the variation, respectively. All groups of geochemical parameters showed significant relationships with bacterial and fungal data (Table 2), the strongest being displayed with siderophile-, biophile-, lithophile major elements and solution parameters (Table 2). Gold-pathfinder and rare earth elements were capable of explaining 24.8% and 42% of the variation, respectively (Table 2).

At Humpback and Wildcat, that is, the contiguous sites at Lawlers, bacterial community structures varied with underlying lithology ($\sqrt{CV}=10.8$, PERMANOVA $P<0.05$) and landform ($\sqrt{CV}=9.7$, PERMANOVA $P<0.1$; Table 1), but no links to the underlying Au deposit were detected (Table 1). Fungal communities varied with lithology, Au deposit and landform ($P<0.1$; Table 1); archaeal communities varied with landform ($P=0.07$; Table 1). Strong separation of bacterial and fungal communities occurred across sites with different lithologies ($P<0.001$) and landforms ($P<0.001$; Figure 2e, Supplementary Figure S2C). Canonical correlations were high ($\delta 1=0.96$; $\delta 2=0.81$), and CA1 and CA2 accounted for 15.9% and 9.6% of variation in the bacterial data set, respectively.

Table 2 DISTLM of bacterial and fungal community diversities, based on M-TRFLP data vs geochemical parameters

Community	Solution (E, C, pH)	Biophile majors (C, N, P, S)	Lithophile majors (Al, Ca, K, Mg, Na)	Siderophile majors (Fe, Mn)	Calcophile minors (Cu, Ga, Ge, Sn, Zn)	Au-pathfinder minors (Ag, As, Au, Bi, Mo, Pb, Se, W)	Lithophile minors (Cr, Nb, Sc, Sn, Sr, Th, U, V)	REEs minor (Ce, Dy, Er, Eu, Gd, Ho, La, Lu, Nd, Pr, Sm, Tb, Tm, Y, Yb)	Siderophile-minor (Co, Ni, Os)
Old Pirate									
Bacteria Pseudo-F	6.1	2.7	2.8	2.6	2.2	1.7	2.1	1.2	1.9
% of variation	18.4/9.2	16.8/4.2	21.5/4.3	8.6/4.3	17.5/3.5	21.6/2.7	25.6/3.2	30.0/2.0	10.2/3.4
Fungi Pseudo-F	3.4	2.4	2.5	2.6	1.8	1.4	1.9	1.2	2.0
% of variation	10.2/5.1	3.6	3.7	4.1	2.8	2.2	2.7	1.9	3.2
Tomakin									
Bacteria Pseudo-F	2.3	2.5	2.3	2.6	1.9	1.7	1.8	1.7	2.2
% of variation	8.8/4.4	18.0/4.5	20.5/4.1	10.0/5.0	18.0/3.6	24.8/3.1	26.4/3.3	42.0/2.8	13.2/4.4
Fungi Pseudo-F	2.3	2.8	3.1	3.7	2.5	1.5	2.3	1.6	3.6
% of variation	8.8/4.4	20.0/5.0	26.5/5.3	13.4/6.7	22.0/4.4	23.2/2.9	31.2/3.9	42.0/2.8	20.1/6.7
Humpback									
Bacteria Pseudo-F	1.7	1.3	1.2	1.3	NS	1.2	NS	NS	NS
% of variation	20.2/10.1	32.8/8.1	39/7.8	17.5/8.5	NS	68.6/7.3	NS	NS	NS
Fungi Pseudo-F	1.8	1.4	1.3	1.6	1.2	1.3	1.4	NS	NS
% of variation	21.4/10.7	34.4/8.6	41.0/8.2	19.4/9.7	37.0/7.4	59.2/7.4	61.6/7.7	NS	28.2/9.4
Wildcat									
Bacteria Pseudo-F	1.5	1.6	1.6	NS	1.4	NS	1.4	1.3	2.1
% of variation	6.2/3.1	12.8/3.2	3.1	NS	2.7	NS	2.8	2.4	12.3/4.1
Fungi Pseudo-F	1.4	1.6	1.3	NS	NS	NS	1.3	1.4	2.2
% of variation	5.8/2.9	13.2/3.3	2.6	NS	NS	NS	2.5	2.5	12.6/4.2

Abbreviations: E.C., electrical conductivity; DISTLM, distance-based linear regression modelling; M-TRFLP, multiplex-terminal restriction fragment length polymorphism; REE, rare earth element. Bold font: significant <0.001 ; normal font: significant <0.05 , italic font: significant <0.1 ; NS, not significant >0.1 .

Because of the lithological differences between Wildcat and Humpback, assessment of associations of community structures with geochemical parameters was conducted separately. At Humpback, major elemental groups and Au-pathfinder elements showed significant relationships with bacterial communities (Table 2). Solution parameters displayed the strongest relationship with bacterial and fungal communities; 60–70% of variation could be explained by Au-pathfinder elements (Table 2). At Wildcat, all major and minor elements, except Au-pathfinder elements, were related to bacterial and fungal communities (Table 2).

M-TRFLP—fingerprinting of microbial communities in microcosms

To assess the effect of soluble Au(III) complexes on community compositions, soil microcosms were amended with up to 100 000 ng Au(III)-chloride g^{-1} soil d.w. Throughout the experiment the number of TRFs in unamended microcosms was 50–60 in Old Pirate soils, but decreased in Wildcat and Humpback microcosm soils (Supplementary Table S4). TRF numbers decreased for bacteria and archaea with increasing concentrations of soluble Au(III) complexes (Supplementary Table S4). In contrast, the numbers of fungal TRFs did not respond to increasing Au concentrations (Supplementary Table S4). Amendment with 100 000 ng Au(III)-chloride g^{-1} soil d.w. led to a reduction of TRFs by >90%; hence these samples were excluded from statistical testing (Supplementary Table S4). The structure of the bacterial community was most strongly characterised by site ($\sqrt{\text{CV}}=21.6$), then Au-amendment ($\sqrt{\text{CV}}=17.4$) and then incubation time ($\sqrt{\text{CV}}=11.2$ (all $P<0.001$; Table 3). Interactive effects were significant, but weaker than the main effects of individual factors (Table 3). Fungal communities

were similarly influenced, whereas the influence of these factors was weaker in archaeal communities (Table 3).

Whereas the amendment with 1000 and 100 000 ng Au(III)-chloride g^{-1} soil d.w. directly affected on the number of TRFs detected (Table 4), the number of TRFs remained approximately constant with loads of 50 ng Au(III)-chloride g^{-1} soil d.w. Yet, the distribution of TRFs, that is, the species composition, varied significantly between amended and unamended samples after 30 days of incubation (Figure 3). In microcosms with Old Pirate samples, canonical correlations were high ($\delta_1=0.99$; $\delta_2=0.96$), and CA1 and CA2 accounted for 32.1% and 21.5% of the variation in the bacterial data set, respectively (Figure 3a). In microcosms with Humpback samples, canonical correlations were high ($\delta_1=0.99$; $\delta_2=0.9$), and CA1 and CA2 accounted for 33.2% and 30% of the variation in the bacterial data set, respectively (Figure 3b). In microcosms with Wildcat samples, canonical correlations were high ($\delta_1=0.99$; $\delta_2=0.95$), and CA1 and CA2 accounted for 38.1% and 25.3% of the variation in the bacterial data set, respectively (Figure 3c).

PhyloChip analyses of bacterial communities

PhyloChip microarrays were used to compare bacterial communities across representative auriferous and non-auriferous samples. Combined across all samples 42 phyla, 77 classes, 146 orders and 222 families were detected. In total, between 957 and 1678 taxa were observed in individual samples, of which 628 were shared across sites (Figure 4a).

To assess which taxa dominated communities, the top 500 probe-intensity score data for each sample were used. *Proteobacteria* (34.4–49.8%), *Firmicutes* (8.4–21.8%), *Actinobacteria* (11.1–23.0%), *Acidobacteria* (1.9–5.8%) and *Cyanobacteria* (0.6–4.8%)

Table 3 Results of PERMANOVA testing Au amendment, IT and field site on the microbial community structure in microcosm experiments

Community structure	Bacteria			Fungi			Archaea		
	CV ($\sqrt{}$)	Pseudo-F	P	CV ($\sqrt{}$)	Pseudo-F	P	CV ($\sqrt{}$)	Pseudo-F	P
Au	17.4	22.6	0.0001 ^e	14.4	6.0	0.0001	18.4	3.9	0.001
SS	21.6	15.0	0.0001	33.9	12.6	0.0001	20.5	3.1	0.007
IT	11.2	7.5	0.0001	11.0	3.1	0.0001	15.7	2.3	0.02
Au \times SS ^a	15.2	6.4	0.0001	18.7	3.8	0.0001	17.6	1.8	0.04
Au \times IT ^b	7.7	2.4	0.0001	4.9	1.2	0.0828	−3.2	1.0	0.5
SS \times IT ^c	9.6	2.6	0.0004	13.5	2.0	0.0001	−7.7	0.9	0.6
Au \times SS \times IT ^d	8.4	1.6	0.0025	13.4	1.5	0.0001		2.4	0.004
Residual	19.4			33.6			43.4		

Abbreviations: AU, gold amendment; IT, incubation time; PERMANOVA, permutational multivariate analysis of variance; SS, sampling site.

^aInteractive effect of Au and SS.

^bInteractive effect of Au and IT.

^cInteractive effect of SS and IT.

^dInteractive effect of Au, SS and IT.

^eSignificance level to assess the criteria was $P<0.05$ (in bold).

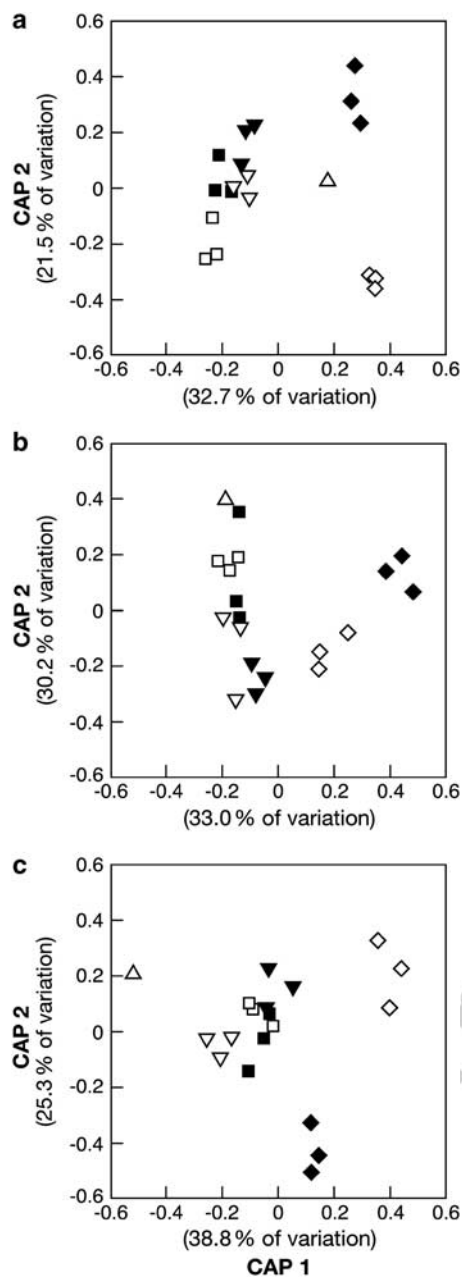


Figure 3 Ordination plots of the first two CAPs produced by CAP comparing bacterial TRLFP patterns from microcosm soils amended with $50 \text{ ng g}^{-1} \text{ AuCl}_4^- (\text{aq})$ (filled symbols) to unamended incubations (unfilled symbols after 0 (Δ), 2 (∇), 10 (\blacksquare , \square) and 30 (\blacklozenge , \lozenge) days of incubation with samples from Old Pirate (a), Humpback (b) and Wildcat (c).

were dominant in all samples (Figure 4b). *Sphingobacteria*, *Verrucomicrobiae*, *Anaerolineae*, *Planctomycetacia*, *Catabacter*, *Bacteroidetes*, *Spirochaetes* and *Chloroflexi* each contributed between 0.6% and 1.8% to the composition of bacterial communities (Figure 4b). Within the *Proteobacteria* class, most taxa were in the *Alpha-proteobacteria* class, consisting of 12–26% of array hits (Figure 4b). At Old Pirate the abundance of *Alpha-proteobacteria*, especially *Sphingomonadales*, and *Firmicutes* (*Bacilli*) was

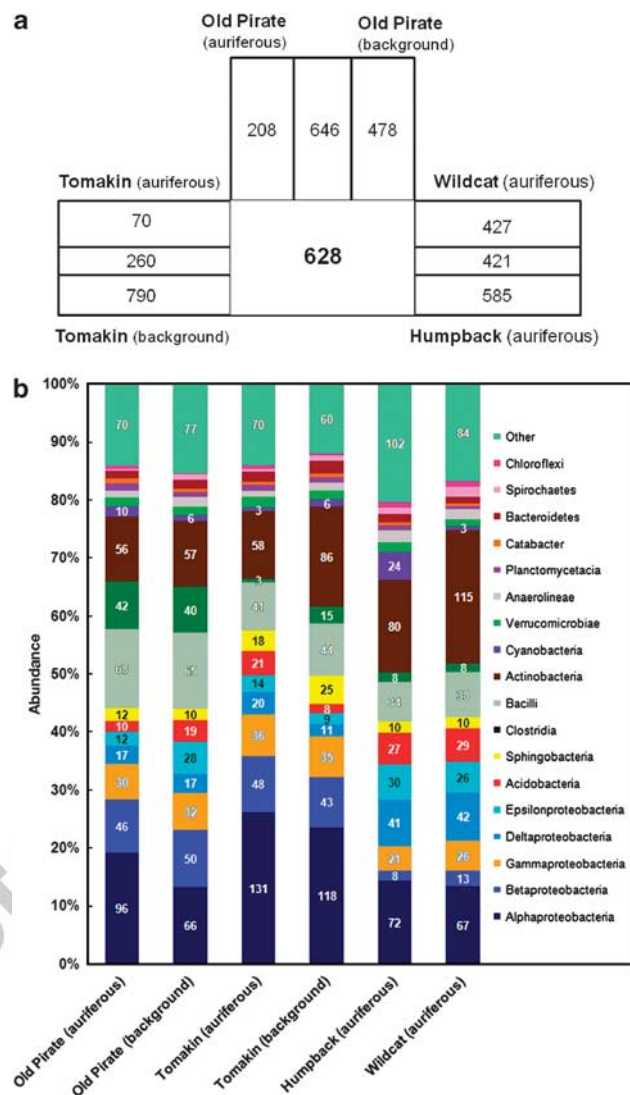


Figure 4 Venn diagram of total detected and shared taxa (a), and distribution of major phyla (b), at the in depth Old Pirate, Tomakin Park, Humpback and Wildcat sampling sites; with classes shown for *Proteobacteria* and *Firmicutes*, phyla with <1% coverage are aggregated into other. PlyloChip data with presence indicated by $\text{pf} > 0.9$.

significantly higher in auriferous soils; these taxa explained >30% of the dissimilarity to non-auriferous soils (Supplementary Table S5). *Beta-* and *Epsilon-proteobacteria* and *Acidobacteria* were less abundant in auriferous soils; differences in these taxa explained 20% of dissimilarities (Supplementary Table S5). Similar results were obtained at Tomakin, where *Actinobacteria*, *Alpha-proteobacteria* and *Bacilli* were more abundant, and *Beta-* and *Epsilon-proteobacteria* and *Acidobacteria* were less abundant in auriferous soils (Supplementary Table S5). At Wildcat and Humpback sites, the strongest effect on microbial communities was lithology; hence two auriferous soils from erosional zones were analysed. At Wildcat, *Actinobacteria*, *Beta-* and *Gamma-proteobacteria* were more

abundant and contributed ~37% to dissimilarity to Humpback soil. Humpback soils were rich in *Cyanobacteria* explaining 18% of dissimilarities (Supplementary Table S5). Collective analysis across all samples showed that a 'core community' of 50 taxa occurred only in auriferous soils (Supplementary Table S6). In all, 13 of these 50 taxa are among the 500 highest ranking taxa. Organisms hybridising to *C. metallidurans* probes occurred in auriferous Tomakin-, Wildcat- and Humpback soils and all Old Pirate soils.

In microcosms incubated for 10 days with 1000 ng g^{-1} d.w. soil of AuCl_4^- , *Actinobacteria* and *Bacilli* were more, *Acidobacteria*, *Clostridia*, *Beta*- and *Gamma*-, *Delta*- and *Epsilonproteobacteria* were less abundant compared with unamended controls (Figure 5a, Supplementary Table S7). In microcosms incubated for 30 days with 50 ng g^{-1} d.w. soil of AuCl_4^- , *Alpha-proteobacteria*, *Actinobacteria* and *Bacilli* were more, and *Acidobacteria* and *Beta-proteobacteria* were less abundant (Figure 5b). The Sphingomonads, a member of the *Alphaproteobacteria*, were enriched in Au-amended samples (Figures 5a and b).

GeoChip-analysis of functional potential

The composition of functional genes associated with geochemical cycling and metal resistance varied between the auriferous and non-auriferous soils from Old Pirate and Tomakin (Figure 6a). Gene families associated with the degradation/transformation of organic contaminants (33.1%) and for

metal resistance/reduction (18.5%) constituted the majority of genes detected (Figure 6b).

At Old Pirate and Tomakin, 1015 and 908 metal-resistance/reduction genes were detected in auriferous soil compared with 872 and 843 in non-auriferous soils, respectively (Figure 6b). In particular, Cu-, Cr- and As-detoxification genes were present in higher abundances (Figure 6b). At Old Pirate, 4 of 10 genes contributing most to the differences ($P < 0.05$) between communities in auriferous and non-auriferous soils were metal-resistance genes: *chrA*, *czcD*, *copA* and *zntA*, which are involved in Cr-, Cd-Zn-Co-, Cu-Zn- transport and detoxification, respectively (Supplementary Table S8). At Tomakin, variation in the abundances of the *copA* gene was most strongly linked to separation between auriferous and non-auriferous soils, (Supplementary Table S9). Across all sites, 233 genes occurred only in auriferous soils. Of these, 69 are metal-resistance/reduction genes, and 19 of these are known to be involved in Cu resistance

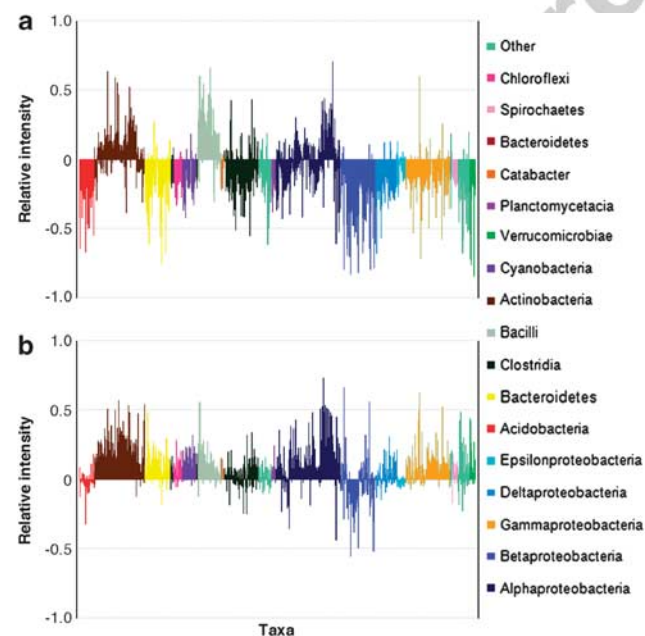


Figure 5 Plots showing relative differences of taxa observed from Phylochip data of Au-amended and -unamended microcosm experiments with soils samples from Old Pirate. (a) Differences observed after 10 days of incubation with 1000 ng g^{-1} AuCl_4^- d.w. soil; (b) Differences observed after 10 days of incubation with 50 ng g^{-1} AuCl_4^- d.w. soil.

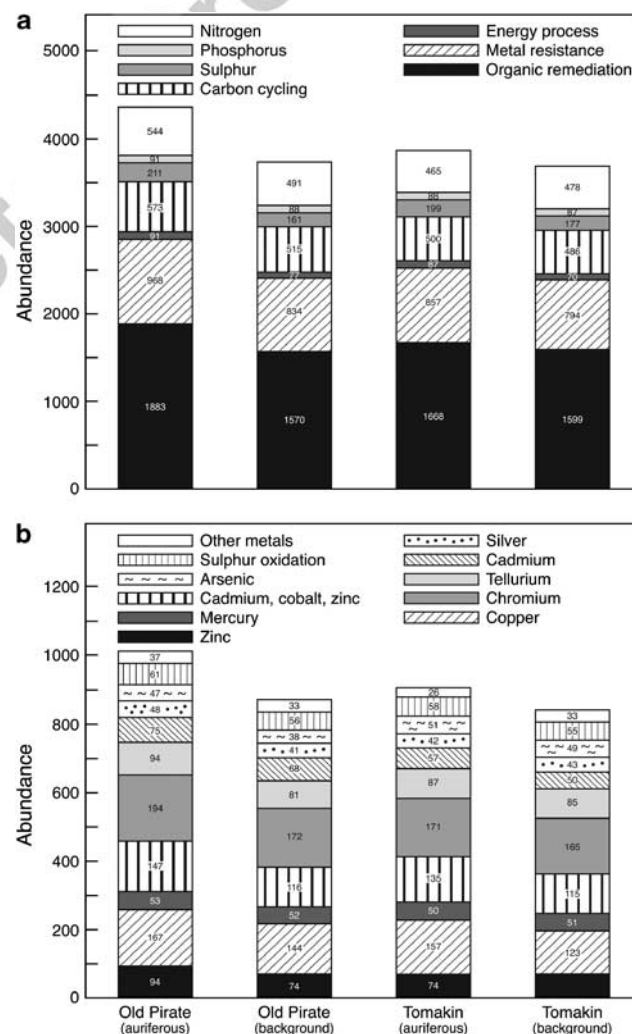


Figure 6 Distribution of major gene categories (a) and metal reduction/detoxification genes (b) in key samples from Old Pirate and Tomakin Park determined by GeoChip analysis.

(Supplementary Table S10). In particular, genes that hybridised with the metal-resistance genes *copA*, *chrA* and *czcA* of the aurophillic bacterium *C. metallidurans* CH34 occurred only in auriferous soils (Supplementary Table S10).

Discussion

Our results showed that geogenic factors and associated geochemical parameters are fundamental determinants of microbial communities in naturally metal-rich soils. In addition to affecting the phylogenetic structure of soil microbial communities, geogenic factors also affected the community at a functional level. Phylogenetically, communities primarily differed across landforms (Old Pirate, Humpback and Wildcat), lithology (Humpback and Wildcat) and with underlying Au deposits (Old Pirate and Tomakin). Lithology and landform have long been recognised as primary drivers of soil formation processes, soil geochemistry, depth of weathering and the biogeographical distribution of plant communities (Jenny, 1941; Viles *et al.*, 2008). It has been hypothesised that these factors also influence microbial community composition and function (for example, Viles, 2011, in press). However, in studies spanning both wide (France and California), and narrow geographic ranges (for example, glacier forefields) no significant links of communities to lithologies and landforms have been detected (Castellanos *et al.*, 2009; Dequiedt *et al.*, 2009; Lazzaro *et al.*, 2009; Drenovsky *et al.*, 2010). In these studies, factors linked to landuse (for example, farming and forestry) and climatic conditions, were dominant. Particularly in Europe this may be the result of permanent human settlement, with its influence on landuse for forestry, intensive agriculture, industry and recreation. In Australia the history of such influences is shorter, and the study sites have not been subjected to intense agriculture or forestry, hence geogenic drivers have not been masked by the shorter-term influences. This was also observed in two other Australian studies, which have shown that soil surface geomorphology and the presence of buried mineralisation influenced resident microbial communities (Mele *et al.*, 2010; Wakelin *et al.*, 2012a). A study by Wagai *et al.* (2011) has shown that soil microbial communities in Bornean tropical forests were strongly influenced by the underlying parent material; here the influence occurred either directly via soil geochemistry or indirectly via floristic expression.

Our study points to a stronger influence of soil metal contents on microbial communities in natural metal-rich soils than previously reported (for example, Duxbury and Bicknell (1983). There may be several explanations for this observation. These are related to the long formation history of metal-rich soils used in this study, which controls the speciation and hence biological activity of

mobile metals, the low organic matter contents and the solution parameters. Geogenic factors have had more time to express their influences on the geochemistry of many Australian landscapes compared with those in Europe or North America. At our study sites, the long weathering history led to the development of deeply weathered profiles and mature soils in landscape settings that are Tertiary to Permian. There the landscapes are comparatively young, with 'new' landforms derived from extensive Upper Pleistocene glaciations overlying 'fresh' rocks.

Owing to the long weathering history metals that are commonly assumed to be immobile, for example, Au, As, Pb, Ag, Cd, Cu, Zn, Ni and U, are highly mobile and hence bioavailable in mature Australian soils, which affects microbial communities (Aspandiar *et al.*, 2008). Studies on soils with high C contents have shown that differences in soil organic matter quality are strong drivers of difference in microbial communities (for example, Torsvik and Øvreås, 2002). This may mask the influences of mobile metals, because soil organic matter displays strong metal sorption capacities, which decreases the mobility and bioavailability of mobile metals (Aspandiar *et al.*, 2008). The low organic matter contents of our soils, and those studied by Wakelin *et al.* (2012a), may have increased metal mobility. At all sites, the influence of soil-pH and E.C. on microbial community compositions was greater than that to metal contents. This shows that solutions, which directly interact with the cell surfaces, have most immediate influence on microbial communities. Solution composition of are the result of interactions between inorganic and organic parent material, biota and water. Solution parameters reflect this and the mobility of metals (Aspandiar *et al.*, 2008): E.C. correlates to the concentrations of major mobile elements, for example, Na, K, Ca, Mg, and S; pH correlates positively to trace metal contents, for example, Au, As and W at Old Pirate, and Cd, Cr, Cu, Ni, Zn and rare earth elements at Humpback. Hence, in the low C soil, mobile metals in solutions, whose composition is controlled by geogenic factors, can have a short-term impact on soil microbial communities, as demonstrated in the Au-amended microcosms.

Communities in auriferous soils contain more unique taxa and greater abundances of metal-resistance genes (Supplementary Table S6; Supplementary Table S10). This suggests that evolutionary pressures, possibly associated with low C content and elevated concentrations of mobile metals were involved in selecting for specialised communities. *Alpha-proteobacteria*, especially *Sphingomonas* spp., *Actinobacteria* and *Bacilli* were enriched in auriferous soils and Au-amended microcosms. In contrast, studies conducted in metal-polluted soils have often found higher abundances of *Beta-proteobacteria*; this was not observed here or by Wakelin *et al.* (2012b).

This may be linked to the low C and N contents of studied soils, as *Alpha-proteobacteria* have a high proportion of genera that are phototrophic and/or are capable of N fixation, making them ideally suited to this low C and N habitat.

Metal-driven selection may explain the higher abundances of *Sphingomonas* spp. and taxa/metal-resistance genes related to *C. metallidurans* in auriferous soil. *Sphingomonas* are Gram-negative aerobic bacteria that are ubiquitous in soils and have a wide physiological tolerance to Cu and other metals (White *et al.*, 1996). The occurrence of taxa related to *C. metallidurans* may be even more specifically linked to elevated concentrations of mobile Au, as *C. metallidurans* reductively precipitates toxic Au(I/III) complexes from solution via Au-regulated gene expression (Reith *et al.*, 2009). Recent studies using transcriptome microarrays have shown that genes in *C. metallidurans* CH34 assumed to be specific to Cu resistance, that is, *cup/cop* (*cupCAR*) and (*copVTMKNRABC-DIJGFLQHE*), were upregulated with Au(III) complexes at concentrations far below those where Cu regulation occurs (Reith *et al.*, 2009). In particular, the Cu-transporting P-type ATPase *copA* was strongly upregulated in the presence of Au(III) complexes, suggesting its involvement in their detoxification (Reith *et al.*, 2009). However, Cu is not enriched in the auriferous soils at Old Pirate or Tomakin (this study; Reith *et al.*, 2005), suggesting an influence of mobile Au complexes on the abundance of *copA* genes.

Conclusions

The microbial ecology of soils is strongly influenced by geogenic factors. Phylogenetic structure, especially bacterial, is strongly linked to lithology, presence underlying deposits and landforms, which control the geochemical make-up these soils. Furthermore, geochemical properties select for communities enriched with the functional potential to deal with elevated concentration of toxic metals. In particular, abundances of *Alpha-proteobacteria*, especially *Sphingomonas* spp. and *Firmicutes* (*Bacillus* spp.) were higher in auriferous compared with non-auriferous soils.

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Q19

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